Simple Decay: Radioactive Parent ⇒ *Stable Daughter*



Decay Series: Radioactive Parent \Rightarrow Radioactive Daughter $N_1(T_{1/2} = 10 \text{ hr})$ Decay of parent Ν $N_2(T_{1/2} = 1 hr)$ Ingrowth from parent and decay of daugther 10 11 12

Image by MIT OpenCourseWare. After Figure 4.4 in Faure.

Time in hours

Consider the decay series $N_1 \rightarrow N_2 \rightarrow N_3$

Remember:

 $-dN_1/dt = \lambda_1 N_1$

Now we consider N_2 , that is produced by decay of N_1 and itself decays to N_3 :

2)
$$dN_2/dt = \lambda_1 N_1 - \lambda_2 N_2$$

Remember:

3)
$$N_1 = N_{1,0} e^{-\lambda l t}$$

Substitute 3) into 2):

4)
$$dN_2/dt = \lambda_1 N_{1,0} e^{-\lambda_1 t} - \lambda_2 N_2$$

Rearrange:

5)
$$dN_2/dt + \lambda_2 N_2 - \lambda_1 N_{1,0} e^{-\lambda_1 t} = 0$$

Solving this first order differential equation for N₂ yields:

6)
$$N_2 = \lambda_1 / (\lambda_2 - \lambda_1) N_{1,0} (e^{-\lambda_1 t} - e^{-\lambda_2 t}) + N_{2,0} e^{-\lambda_2 t}$$

The solution, as well as equivalent solutions for three nuclides and the general case, are known as Bateman (1910) equations/solutions.

The first term in equation 6) is the number of N_2 atoms decayed from N_1 not yet decayed to N_3 The second term in equation 6) is number of N_2 atoms that remain from the initial $N_{2,0}$

The classic Bateman paper on the famous "Bateman equations"

Bateman, Harry. "The solution of a system of differential equations occurring in the theory of radioactive transformations." In *Proc. Cambridge Philos. Soc*, vol. 15, no. pt V, pp. 423-427. 1910.

The solution of a system of differential equations occurring in the theory of radio-active transformations. By H. BATEMAN, M.A., Trinity College.

[Read 21 February 1910.]

1. It has been shown by Prof. Rutherford * that the amounts of the primary substance and the different products in a given quantity of radio-active matter vary according to the system of differential equations,



where P, Q, R, S, T, \dots denote the number of atoms of the primary substance and successive products which are present at time t.

Prof. Rutherford has worked out the various cases in which there are only two products in addition to the primary substance, and it looks at first sight as if the results may be extended to any number of products without much labour.

Unfortunately the straightforward method is unsymmetrical and laborious, and as the results of the calculations are needed in some of the researches which are being carried on in radio-activity the author has thought it worth while to publish a simple and symmetrical method of obtaining the required formulae.

2. Let us introduce a set of auxiliary quantities p(x), q(x), r(x), ... depending on a variable x and connected with the quantities P(t), Q(t), R(t), ... by the equations,

$$p(x) = \int_0^\infty e^{-xt} P(t) dt, \quad q(x) = \int_0^\infty e^{-xt} Q(t) dt....(2).$$

It is easily seen that

$$\int_{0}^{\infty} e^{-xt} \frac{dP}{dt} dt = -P(0) + x \int_{0}^{\infty} e^{-xt} P(t) dt \dots \dots \dots (3),$$

= -P_0 + xp,

* Radio-activity, 2nd edition, p. 331.

continue...

If there are no atoms of the original daughter $N_{2,0}$, then 6) simplifies to:

7)
$$N_{2} = \lambda_{1} / (\lambda_{2} - \lambda_{1}) N_{1,0} (e^{-\lambda t} - e^{-\lambda 2t})$$
Example: ²³⁸U \rightarrow ²³⁴U \rightarrow ²³⁰Th \rightarrow ... \rightarrow ²⁰⁶Pb

8) $(^{230}\text{Th}) = (^{230}\text{Th})_x + (^{230}\text{Th})_s$ x = excess, not supported, s = supported from ^{238}U

Let's first consider the excess activity only, at some time, t:

9)
$$(^{230}\text{Th})_x = (^{230}\text{Th})_{x,0} e^{-\lambda 230}$$

Normalize by a "stable" isotope. Relative to the short-lived daughters, ²³²Th is "stable"

10)
$$({}^{230}\text{Th}/{}^{232}\text{Th})_x = ({}^{230}\text{Th}/{}^{232}\text{Th})_{x,0} e^{-\lambda 230 t}$$

Now let's consider the 238 U-supported (230 Th) – see equation 7)

11)
$$({}^{230}\text{Th})_{s} = \lambda_{234}/(\lambda_{230} - \lambda_{234}) {}^{234}\text{U}_{0} (e^{-\lambda_{234}t} - e^{-\lambda_{230}t})$$

At secular equilibrium $(^{234}U) = (^{238}U)$ then $^{234}U \lambda_{234} = (^{238}U)$

Also $\lambda_{230} - \lambda_{234} \cong \lambda_{230}$ as λ_{234} is small: $e^{-\lambda 234 t} \cong 1$

²³⁸U Decay Series



Peucker-Ehrenbrink, 2012

²³⁸U Decay Series



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Application in determining chronologies of sediments

Basic decay equation: $N = N_0 e^{-\lambda t}$ assuming no U-supported activity Replace time (t) with depth in sediment column (d) divided by sedimentation rate (sr)

t = d / sr

Decay equation: $N = N_0 e^{-\lambda d/sr}$ ln N = ln N₀ - $\lambda d/sr$

t = d / sr

In a diagram of ln N (y axis) and d (x axis) the slope (m) is

Figure of $log(^{230}Th/^{232}Th)_A$ versus Core Depth, cm, has been removed due to copyright restrictions.

Faure, Gunter. "Principles of Isotope Geology." (John Wiley & Sons, 1986): 367. $m = -\lambda / sr$

and

$$sr = -\lambda / m$$

This method is also known as the ionium (²³⁰Th) method of dating.

Complications...



Depth in the core

Complications...



continue...

If there are no atoms of the original daughter $N_{2,0}$, then 6) simplifies to:

7) $N_2 = \lambda_1 / (\lambda_1 - \lambda_2) N_{1,0} (e^{-\lambda_1 t} - e^{-\lambda_2 t})$

Example: ${}^{238}U \rightarrow {}^{234}U \rightarrow {}^{230}Th \rightarrow \dots \rightarrow {}^{206}Pb$

8) $(^{230}\text{Th}) = (^{230}\text{Th})_x + (^{230}\text{Th})_s$ x = excess, not supported, s = supported from ^{238}U

Let's first consider the excess activity only, at some time, t:

9)
$$({}^{230}\text{Th})_{x} = ({}^{230}\text{Th})_{x,0} e^{-\lambda 230 t}$$

Normalize by a "stable" isotope. Relative to the short-lived daughters, ²³²Th is "stable"

10)
$$({}^{230}\text{Th}/{}^{232}\text{Th})_x = ({}^{230}\text{Th}/{}^{232}\text{Th})_{x,0} e^{-\lambda 230 t}$$

Application!

Now let's consider the 238 U-supported (230 Th) – see equation 7)

11)
$$({}^{230}\text{Th})_{s} = \lambda_{234}/(\lambda_{230} - \lambda_{234}) {}^{234}\text{U}_{0} (e^{-\lambda 234t} - e^{-\lambda 230t})$$

At secular equilibrium, $(^{234}U) = (^{238}U)$ then $^{234}U \lambda_{234} = (^{238}U)$

Also
$$\lambda_{230} - \lambda_{234} \cong \lambda_{230}$$
 as λ_{234} is small: $e^{-\lambda 234 t} \cong 1$

continue...

11)

$${}^{230}\text{Th}_{s} = \lambda_{234} / (\lambda_{230} - \lambda_{234}) {}^{234}\text{U}_{0} (e^{-\lambda 234t} - e^{-\lambda 230t})$$

$$\lambda_{230} {}^{230}\text{Th}_{s} = \lambda_{234} {}^{234}\text{U}_{0} (1 - e^{-\lambda 230t})$$
12)

$$({}^{230}\text{Th}_{s} = {}^{238}\text{U}_{0} (1 - e^{-\lambda 230t})$$

From equation 9) $(^{230}\text{Th})_x = (^{230}\text{Th})_{x,0} e^{-\lambda 230 t}$

$$(^{230}\text{Th}) = (^{230}\text{Th})_{x,0} + (^{230}\text{Th})_{s}$$

Total 230 Th activity = initial excess 230 Th activity + 238 U-supported activity

13)
$$(^{230}\text{Th}) = (^{230}\text{Th})_{x,0} e^{-\lambda 230t} + ^{238}\text{U}(1 - e^{-\lambda 230t})$$

Normalize to ²³²Th

14)
$$(^{230}\text{Th}/^{232}\text{Th}) = (^{230}\text{Th}/^{232}\text{Th})_{x,0} e^{-\lambda 230t} + (^{238}\text{U}/^{232}\text{Th}) (1 - e^{-\lambda 230t})$$

If (²³⁰Th/²³²Th) is plotted against (²³⁸U/²³²Th), equation 14) is a linear equation, the so-called

²³⁰Th-²³⁸U isochron diagram

²³⁰Th-²³⁸U isochron diagram



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Modified after figure 21.7 of Faure, Gunter. "Principles of Isotope Geology." (John Wiley & Sons, 1986): 379.

²³⁸U-series



²³⁵U-series

235U

0.704 by

²³¹Th

1.06 d





Peucker-Ehrenbrink, 2012



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²¹²P0

3 10⁻⁷ s

²⁰⁸Pb

stable

1.01 h

²⁰⁸TI

3.05 m

²²⁶Ra - ²³⁰Th disequilibrium diagram

Figure of radioactive disequilibria between ²²⁶Ra-²³⁰Th and ²³⁸U-²³⁰Th in lavas from austral and southern volcanic zones removed due to copyright restriction.

After Figure 1 in Sigmarsson, O., J. Chmeleff, J. Morris, and L. Lopez-Escobar. "Origin of ²²⁶ Ra–²³⁰Th disequilibria in arc lavas from southern Chile and implications for magma transfer time." *Earth and Planetary Science Letters* 196, no. 3 (2002): 189-196.

²³⁰Th-²³⁴U-²³⁸U activity ratio diagram



Courtesy of Mineralogical Society of America. Used with permission.

From: Chabaux, Riotte and Dequincey, U-Th-Ra Fractionation during weathering and river transport. In: Bourdon et al. (Eds), Uranium-series geochemistry, Rev. Mineral. Geochem., vol. 52, chapter 13, 533-576.



Peucker-Ehrenbrink, 2012

Further Reading...

•Bourdon, B. et al. (Ed.) Uranium-Series Geochemistry, Rev. Mineral. & Geochem., vol. 52, p. 20-21 (general solution to U-series equations).

•Bourdon, B. et al. (Ed.) Uranium-Series Geochemistry, Rev. Mineral. & Geochem., vol. 52, p. 2-7 (U-series introduction).

•DecaySeries toy (simple Excel file).



Measurement Uncertainties

All measurements are afflicted with uncertainties. For large number of events, binomial distributions asymptotically approach Gaussian (or normal) distributions. The spread in events (here numerical values of isotope ratios, count rates or ion currents) is equal to \sqrt{N} . According to Gaussian statistics about 2/3 of the results lie within the range N $\pm \sqrt{N}$ (one standard deviation), about 95% lie within the range N $\pm 2\sqrt{N}$ (two standard deviations), and ~99% lie within the range N $\pm 3\sqrt{N}$. The fractional uncertainty is thus $\sqrt{N/N}$, or $1/\sqrt{N}$. If you measure twice as long (N*) you get twice as many events

$$N^* = 2N$$

the fractional uncertainty is
$$\sqrt{(2N)/2N} = 1/\sqrt{(2N)}$$

i.e. $= \sqrt{1/2 * 1/\sqrt{N}}$

reducing the fractional uncertainty only by $\sim 30\%$. The fractional uncertainty improves only as the square root of time (or ion current, or count rate). If you attempt to improve the uncertainty by a factor of two, you need to measure four times as long, or measure a four-times stronger ion current.

In order to evaluate if uncertainties associated with small ion beam intensities significantly affect the measured ratios it is often helpful to assume that all uncertainties are associated with uncertainties in the smallest ion current (least abundant isotope). By assuming an arbitrary uncertainty in the measurement of this ion current you can plot an error trend on plots of isotope ratio versus another isotope ratio (same isotope in the denominator, i.e. $m_2 = m_4$, if m_1 and m_3 are isotopes in the numerator). This trend is often distinct from a instrumental fractionation trend and helps to assess what process dominates the uncertainty of your analysis.

Measurement Uncertainty



The Reporting of Data & Uncertainties

mean concentration
$$= \overline{C} = \sum_{i=1}^{N} \frac{C_i}{N}$$

r.m.s. deviation from average $= \sigma = \left\{ \sum_{i=1}^{N} \frac{(C_i - \overline{C})^2}{N} \right\}^{1/2}$

Text removed due to copyright restrictions.

See page 70 of Warren, Warren S. *The physical basis of chemistry*. Academic Press, 2000.

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