12.335/12.835 EXPERIMENTAL ATMOSPHERIC CHEMISTRY, FALL 2014

TOPIC 1 ATMOSPHERIC PHOTOCHEMISTRY and AIR POLLUTION

INTRODUCTION TO THE CHEMISTRY OF AIR POLLUTION

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OZONE CYCLE & PHOTOCHEMICAL STEADY STATE APPROXIMATION (PSSA)

Derivation of Ozone PSSA



$$\frac{d[O(^{3}P)]}{dt} \approx 0 = k_{1}[NO_{2}] - k_{2}[O(^{3}P)][O_{2}][M]$$

$$k_{1}[NO_{2}] = k_{2}[O(^{3}P)][O_{2}][M]$$

$$\frac{d[O_{2}]}{dt} \approx 0 = -k_{2}[O(^{3}P)][O_{2}][M] + k_{3}[NO][O_{3}]$$

$$k_{3}[NO][O_{3}] = k_{2}[O(^{3}P)][O_{2}][M]$$

$$\boxed{[O_{3}] = \frac{k_{1}[NO_{2}]}{k_{3}[NO]}}$$
The equation of the equation





Fate of NO_x in the Atmosphere



PAN can travel long distances and then serve a downwind source of NO_x

Ozone Isopleth





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The hydroxyl free radical (OH) is the major oxidizing chemical in the global atmosphere

It annually removes about 3.7 billion metric tons (Pg) of trace gases (CO, CH_4 , higher hydrocarbons, hydro-halocarbons, NO_x , SO_x , etc.) from the atmosphere..

With a lifetime of only about 1 second it is possible to measure locally, but not possible to measure directly at regional to global scales.

Use measurements of the industrial chemical CH₃CCI₃, whose emissions are are known and whose major sink is OH, to indirectly estimate large scale OH variations.

Oxidant Sources in the Atmosphere

Hydroxyl Radical (OH)

 Photodissociation of O₃ O₃ + hv → O₂ + O(¹D) O(¹D) + H₂O → 2 OH·
 Photodissociation of HONO and H₂O₂

 $HONO + hv \rightarrow OH \cdot + NO$ $H_2O_2 + hv \rightarrow OH \cdot + OH \cdot$

 Alkene oxidation R₁CH=CHR₂ + O₃ → OH· + prod.

From HO₂·

 $- \text{NO} + \text{HO}_2 \cdot \rightarrow \text{NO}_2 + \text{OH} \cdot$

$H \cdot + O_2 + M \rightarrow HO_2 \cdot + M$ $HCO \cdot + O_2 \rightarrow HO_2 \cdot + CO$ • From alkoxy radical reactions

Hydroperoxyl Radical (HO₂)

 $CH_2O + hv \rightarrow H \cdot + HCO \cdot$

Photodissociation of CH₂O

$$\text{RCH}_2\text{O} \cdot + \text{O}_2 \rightarrow \text{RCHO} + \text{HO}$$

 Byproduct of oxidation of organics

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OH Production Steady State

Production of OH is by reaction of water vapor with O(1D)

(R1) $O_3 + h\nu \rightarrow O_2 + O(^1D)$ (R2) $O(^1D) + M \rightarrow M + O(^3P)$ (R3) $O(^1D) + H_2O \rightarrow 2OH$

Assuming steady state for O(1D) (since R2>>R3)

$$\begin{split} P_{OH} &= 2k_3[O(^1D)][H_2O] = \frac{2k_1k_3}{k_2[M] + k_3[H_2O]}[O_3][H_2O] \\ &\approx \frac{2k_1k_3}{k_2[M]}[O_3][H_2O] \end{split}$$

HOWEVER, IN POLLUTED AIR THERE ARE MULTIPLE SOURCES OF OH DEPENDING ON TIME OF DAY

 Early Morning Source of OH in Polluted Air

 HONO(g) + hv \longrightarrow $\dot{O}H(g) + \dot{N}O(g)$ $\lambda < 400 \text{ nm}$

 Nitrous
 Hydroxyl Nitric

 acid
 radical
 oxide

Mid-Morning Source of OH in Polluted Air

HCHO(g) + hv> Formal- dehyde	HCO(g) + H(g) Formyl Atomic radical hydrogen	λ < 334 nm
H(g) + O ₂ (g) → Atomic Molecular hydrogen oxygen	HO ₂ (g) Hydroperoxy radical	
HĊO(g) + O ₂ (g)	CO(g) + HO ₂ (g) Carbon Hydroperoxy monoxide radical	
NO(g) + HO₂(g) → Nitric Hydroperoxy oxide radical	NO2(g) + OH(g) Nitrogen Hydroxyl dioxide radical	
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Afternoon Source of OH in Polluted Air



BECAUSE OZONE PRODUCTION TAKES TIME, TRANSPORT PLAYS AN IMPORTANT ROLE IN DETERMINING EXPOSURE LEVELS



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USA TROPOSPHERIC COLUMN OZONE BY SEASON (Satellite observations in DOBSON UNITS = 2.7 x 10¹⁶ molecules/cm²)





Image courtesy of Fishman, J., A. E. Wozniak, and J. K. Creilson. From *Atmospheric Chemistry and Physics* 3 (2003): 893–907.

As we will discuss in a later lecture, human health is affected by exposure to ozone expressed here as a function of the AOT40 Index (AOT40 = hourly ozone exposure above 40 ppb in units of ppb.hr/mo.)



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ATMOSPHERIC AEROSOLS

DEFINITION: ALL SUSPENDED LIQUID, SOLID & MIXED LIQUID-SOLID PARTICLES IN AIR EXCEPT WATER DROPLETS & ICE CRYSTALS e.g. BLACK CARBON, SULFATE, ETC.

FIGURES: Aerosol optical depth, Taer at 0.55 µm (color bar) as determined by the MODIS satellite instrument for the **January to March 2001 mean** (top panel) and for August to **October 2001 mean (bottom** panel). The top panel also shows the location of AERONET sites (white squares) that have been operated (not necessary continuously) since 1996. The bottom panel also shows the location of different aerosol lidar networks (red = EARLINET, orange = ADNET, black = **MPLNET).** (Ref: IPCC AR4 WG1, Chap. 2, Fig. 2.11, 2007)

TYPICAL URBAN AEROSOL COMPOSITION, ORIGIN (primary, secondary, either) & SHAPE Sulfates $[H_2SO_4 (aq), NH_4)_2SO_4 (aq, s)]$ Black Carbon [C) (s)]Organic Carbon $[C_xH_yO_z (l, s)]$ Dust [Silicates (s), Clays (s), Pollens (s)] Nitrates $[HNO_3 (aq), NH_4NO_3) (aq, s)]$ Chlorides [NaCl (aq, s)]Mixtures $[(H_2SO_4 \text{ on BC, OC, dust}]$ Spheres (liquids) Crystals (ice, salts, minerals) Dendriform (snow, soot)

TYPICAL URBAN AEROSOL NUMBER (N) DENSITY (n_n), & SURFACE AREA (S) & VOLUME (V) WEIGHTED DENSITIES (n_S, n_V), AS FUNCTIONS OF PARTICLE DIAMETER (D_p)



$$n_N(D_p) = \frac{dN}{dD_p}$$

$$n_S(D_p) = \pi D_p^2 n_N(D_p)$$

$$n_V(D_p) = \frac{\pi}{6} D_p^3 n_N(D_p)$$

Examining the chemistry and physics of air pollution, we see that air pollution regulations affect climate by increasing CH_4 (warming), decreasing H_2SO_4 aerosols (warming), decreasing O_3 (cooling) and decreasing black carbon aerosols (cooling).



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